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²¹⁰Pb-derived sediment accumulation rates across the Wider Caribbean Region

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ABSTRACT

The Wider Caribbean Region is an important tourist destination where agricultural, industrial and shipping activities are also carried on. Coastal zones are heavily populated and receive a high human pressure; however, few monitoring programmes allow assessing long-term anthropogenic impact trends in these areas, which are especially useful for integrated management programs. Through the support of the International Atomic Energy Agency (project RLA/7/012), sediment core activities of ²¹⁰Pb and ¹³⁷Cs were used to evaluate changes in sedimentation rates in 11 relevant coastal areas of the region, where environmental information is scarce, but needed to support national environmental policies. Most ²¹⁰Pb activity profiles were atypical, attributed to non-steady sedimentation conditions; whereas ¹³⁷Cs activity profiles, showing very low values, were of little help for ²¹⁰Pb-dating corroboration. Results evidenced conspicuous changes in mass accumulation rates (MAR), specially through the Anthropocene (i.e. since ~1950s) in most cases, attributed to deforestation and land erosion (one of the clearest indicators of global change), and the input of urban and industrial untreated wastes. The recent MAR decrease in Havana Bay (Cuba) was attributed to the implementation of environmental policies, which showed

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that ²¹⁰Pb-derived reconstruction of environmental changes is also useful to verify the effectiveness of management programs to control land-derived erosion and siltation issues. Since siltation can be detrimental to valuable coastal resources around the world, retrospective evaluations of sedimentation rates, based on ²¹⁰Pbdated cores, are highly recommended to assist integrated coastal zone management programs in the region and elsewhere.

1. Introduction

In order to satisfy the needs of natural resources for a rapidly growing human population, the structure and functioning of the Earth system has undergone relevant changes, identified as a whole as *global change*. Among its most conspicuous effects are the loss of vegetation cover and increased continental erosion due to land-use changes (Vitousek, 1994). The perturbations caused by global change are so important that the recognition of the *Anthropocene* as a new geological time unit, starting approximately in the mid 20th century, is under evaluation (AWG, 2019).

Coastal areas, the natural interface between river basins and oceans, are valuable ecosystems but also among the most strongly impacted by global change, because of the large population (more than 2.8 billion people; >40% of the world's population) living within 100 km of the coast (IOC/UNESCO, IMO, FAO, UNDP, 2011). Watershed deforestation is mostly caused by the change in the traditional uses of land (land-use-change, LUC), increasing the sediment flux (causing siltation) and associated particle-reactive contaminants to coastal zones.

The marine environment is the most important natural resource for the countries in the Wider Caribbean Region (WCR), since approximately 60% of the regional Gross National Product depends on its exploitation (e.g. tourism, fisheries, maritime trade). As population and economies expand, the coastal areas are under increasing human pressure owing to land-based sources and activities, importantly LUC due to urbanization and expansion of the agricultural frontier (associated with unsustainable agricultural and forestry practices) and contamination derived from domestic, industrial and agricultural wastes (PNUMA, 2020). However, little information on the long-term effects of coastal siltation in the region is available. Environmental time series are still scarce, short and mostly present in the developed world. To adapt and hopefully mitigate global change, we need to understand its impacts and their temporal evolution during the industrial era, or at least since 1950s, when global change accelerated, especially in the developing world.

The attribution of sedimentary geochemical signals to past events requires a reliable time-frame, that can only be achieved during the Anthropocene through ²¹⁰Pb dating. Examples of the use of ²¹⁰Pb dating to study global change in the coastal zone include the reconstruction of temporal trends of erosion and sediment accumulation due to LUC, contamination by metals and organic compounds, and recent climate change (e.g. Ruiz-Fernández et al., 2012; Ruiz-Fernández et al., 2014; Ruiz-Fernández et al., 2016; Cuellar-Martínez et al., 2017).

With the support of the International Atomic Energy Agency (Technical Cooperation project RLA/7/012, "Use of Nuclear Techniques to Address the Management Problems of Coastal Zones in the Caribbean Region"), laboratories of the Marine-Coastal Research Network (Red de Investigación Marino-Costera; REMARCO, 2020) collected and analyzed sediment cores from relevant coastal environments in 11 countries of the WCR, namely Colombia, Cuba, Dominican Republic, Guatemala, Haiti, Honduras, Jamaica, Mexico, Nicaragua, Panama and Venezuela. In this work, we evaluate the temporal trends of ²¹⁰Pb-derived sediment accumulation rates (SAR) across the WCR, under the hypothesis that the most conspicuous changes in sediment delivery to coastal environments are related to inland human activities throughout the Anthropocene. The conclusions derived from this study might be limited by i) the representativeness of a single core to reconstruct environmental changes at each ecosystem, and ii) problems with the corroboration of

the ²¹⁰Pb chronologies. However, the general patterns observed among the cores and their implications, highlight the need to protect the health of these valuable ecosystems around the WCR region. The information provided is intended to be useful to decision makers to implement national and regional strategies for sustainable coastal zone management.

2. Study area

The IAEA RLA/7/012 project addressed the study of relevant coastal areas, selected by the national counterparts (OIEA, 2020a), where environmental information was scarce but needed to support management plans. The study sites are (Fig. 1):

2.1. Cartagena Bay (Colombia, COL)

Cartagena Bay is located on Colombia's Caribbean coast, between latitudes 10–26°N and 10–16°N, and between longitudes 75-30°W and 75-35°W. It is a semi-closed bay with an average depth of 16 m (Molares and Mestres, 2012). The Dique channel, an artificial arm of the Magdalena river, built in 18th century, discharges freshwater and a large amount of suspended sediments to the bay, transforming it in a heavily stratified estuary (Pagliardini et al., 1982) in which seasonal changes of salinity, dissolved oxygen concentration and other physical, chemical and biological parameters occur (Molares and Mestres, 2012). Cartagena Bay is considered to be one of the Caribbean's hot spots of pollution (Tosic et al., 2019). The input of wastewaters from human settlements (~880,000 inhabitants) and industrial areas of Cartagena (the "Mamonal"), as well as the fluvial discharge of the Magdalena River, which transports eroded soils from its catchment, are considered to be the most important contamination sources to the bay.

2.2. Havana Bay (Cuba, CUB)

Havana Bay (NW Cuba) is a typical enclosed bay, with a mean depth of 10 m, an area of 5.2 km² and a water mean residence time of 7–9 days. It receives the fluvial discharges of the Luyano and Martin Perez rivers, and the Tadeo, Matadero, Agua Dulce and San Nicolas streams, as well as untreated effluents from numerous sources. It is adjacent to Havana City (2,130,081 inhabitants in 2016; ONEI, 2017) where industrial, commercial and harbor activities have significantly increased since the middle 19th century, with the onset of industrial activities (e.g. oil refining, electric power and gas production). The area of Havana City experienced a fast economic growth during the 20th century, with a high diversity of industries and commercial activities, and a rapid population growth that required massive urbanization (from 250,000 inhabitants in 1899 to 2.2 million inhabitants in 2001). Agriculture and intensive forest exploitation in the bay catchment (about 68 km²) has increased soil erosion and, therefore, sediment input to the bay. This ecosystem is considered a hot spot for many contaminants (Díaz-Asencio et al., 2011), causing severe damage to natural resources, owing to the lack of waste treatment facilities (IIT, 1985; Colantonio and Potter, 2006; Valdes--Mujica, 2008; Armenteros et al., 2009). For its location in the Caribbean Region, the area is frequently affected by storms and hurricanes.

2.3. Haina River Estuary (Dominican Republic, DOM)

Haina River (\sim 79 km long, mean annual flow 9.6 m³ s⁻¹, 564 km² catchment area) discharges to the Caribbean Sea through the Haina port,

to the west of Santo Domingo (2.9 million inhabitants). Haina port (depth range between ~6 and 11 m; HIT, 2020) was built in 1950s and currently is the primary container port for DOM, the primary entrance point for petroleum products to the country, and an important bulk port of agricultural products (Boske et al., 2001). Around 43% of the Haina River catchment is dedicated to grazing and agriculture (main crops include cocoa, citrus, coffee and coconut) (MMARN, 2020). The waters at the highest part of Haina River basin are considered of good quality, but downstream the river is contaminated by residues from cattle ranching and agriculture, by solid wastes and sewage from human settlements, by industrial wastewaters (e.g. paper mill, sugarcane mill, beer factories, power plants) and by port activities (Gómez-Mena et al., 2008). These contaminated waters make their way to the coastal zone, where coral reefs degradation and loss have been attributed to the turbidity and sedimentation increase promoted by the Haina River (Herrera-Moreno et al., 2009). Located at the Haina River mouth, Haina port suffers continuous silting caused by catchement deforestation (Boske, 2001).

2.4. Amatique Bay (Guatemala; GUA)

Amatique Bay is a shallow semi-enclosed water body (140 km long littoral, \sim 560 km² area and 10 m mean depth; Yañez-Arancibia et al., 1999), connected to the low-lying Lake Izabal and to the Mesoamerican Reef System. It receives fluvial inputs of the rivers Motagua, Sarstún and Dulce, and hosts the ports of Barrios, Santo Tomás de Castilla and Livingston. Red mangrove forests surround the eastern side of Amatique Bay, although large sections of these mangroves were destroyed by Hurricane Mitch in 1998 (Fonseca and Arrivillaga, 2003). The main economic activity is fishing (mainly along the eastern bay) including species of economic importance (e.g. shrimp, squid and demersal fish). It is considered to be the most important estuarine ecosystem in Guatemala and makes part of the Punta de Manabique Ramsar site (# 1016, since 2000). Among the main threats to Amatique Bay and its associated ecosystems are the expansion of grazing, agriculture and tourism activities, as well as contamination resulting from inland

erosion, which delivers sediments, agrochemicals and sewage (Yáñez-Arancibia et al., 1999; Andrade et al., 2015; RSIS, 2020b), and hydrocarbon spills owing to shipping activities.

2.5. Port-au-Prince Bay (Haiti, HAI)

Port-au-Prince Bay is a natural harbor adjacent to Port-au-Prince city (2.143 million inhabitants in 2010; ERH, 2019), which hosts ~25% of the total country population. The first port installations were built in 1911 (Noel, 1999) and has a maximum depth of \sim 20 m. The bay is contaminated by household wastes, sludge from pit latrine and sewage, transported by surface runoff from the city watersheds (Emmanuel and Azaël, 1998). High concentrations of heavy metals have been detected in effluents draining from paint manufacturing facilities and hospitals to the bay (Angerville et al., 2005; Lefranc, 2016). In addition, it receives effluents from the Froide and Gris rivers, the latter carrying waste from distilleries, breweries, automobile workshops and a large agricultural area (bananas, sugar and vegetable crops). The bay is also affected by siltation, promoted by inland soil erosion owing to acute deforestation and vegetation cover loss (UNEP, 1999), which started from the beginning of the 19th century, since the country had to sell huge quantities of wood, especially mahogany, to pay a debt acquired in exchange of its independence. Deforestation continues today, since nearly 80% of the population uses wood as a source of energy (UNEP, 2010).

2.6. Puerto Cortés Bay (Honduras, HON)

Puerto Cortés Bay hosts the Central America's largest port, located at a deep natural harbor (maximum depth \sim 14 m), adjacent to the coastal city of Puerto Cortés, considered a growth pole in Honduras. The population of the city is \sim 132,000 inhabitants and it is rapidly growing (INE, 2018). The main sources of pollution are the manufacturing industry, agriculture and commercial activities, including the input of oil residues generated by the port activities; contaminated soils, eroded as a result of LUC for agriculture and cattle raising; and domestic and industrial wastewaters through the rivers Tulián, Medina and



Fig. 1. Sediment core collection sites in the Wider Caribbean Region. The site and sampling coordinates (latitude, longitude; in decimal degrees) where each core was collected are in parenthesis. COL (Cartagena Bay, Colombia; 10.36°N, -75.52°W), CUB (Havana Bay, Cuba; 23.14°N, -82.33°W), DOM (Haina River Estuary, Dominican Republic; 18.41°N, -70.02°W), GUA (Amatique Bay, Guatemala; 15.90°N, -88.71°W), HAI (Port-au-Prince Bay, Haiti; 18.55°N, -72.37°W), HON (Puerto Cortés Bay, Honduras; 15.81°N, -87.97°W), JAM (Kingston Harbour, Jamaica; 17.96°N, -76.74°W), MEX (Coatzacoalcos River Estuary, Mexico; 18.24°N, -94.45°W), NIC (Bluefields Bay, Nicaragua; 11.88°N, -83.75°W), PAN (Almirante Bay, Panama; 9.28°N, -82.38°W), VEN (Cariaco Gulf, Venezuela; 10.48°N, -63.72°W).

Chamelecón.

2.7. Kingston Harbour (Jamaica, JAM)

Kingston Harbour is a bar-built estuary, which serves as the major port of Jamaica. The harbour is an elongated bay, extending 26 km of navigable water from east to west and 6.5 km from north to south, with a total surface area of approximately 51 km² (Wade, 1976) and with depths of up to 18 m. It is bordered by the city of Kingston (669,773 inhabitants in 2018; STATINJA, 2020); over 60,000 people live within 1 km of the harbour's shore (Webber et al., 2019) and sewage from 25% of the population is discharged into the harbor, with limited or no treatment. Water circulation is influenced by low tidal action, wind and freshwater inflows from major gullies and waterways. It has long flushing times, e.g. ~18 days for surface waters, and of several months at the inner harbor (Webber and Webber, 2003). The shape and configuration morphology of the bay has been strongly modified since middle 19th century. Its major fluvial inputs are Rio Cobre, Duhaney River and the Sandy Gully surface water drainage system. Kingston Harbour was reported to be eutrophic as far back as 1968, owing to the release of domestic and industrial wastes from the city of Kingston (Dunbar and Webber, 2003), and contaminated by massive soil erosion and run-off of agrochemicals from the largest and richest agricultural valleys across the Rio Cobre basin (Mansingh and Wilson, 1995).

2.8. Coatzacoalcos River Estuary (Mexico, MEX)

The Coatzacoalcos River is Mexico's third most important river by flow rate (28,679 hm³ year⁻¹). Its lower basin hosts the industrial area of Minatitlan–Coatzacoalcos (one of the largest oil and petrochemical production centers of Mexico), the cargo port of Coatzacoalcos and mineral exploitation areas (e.g. salt, silica sands and sulfur). The region is a critical zone for ecologic regulation and environment protection since early 1990s, owing to the economic importance and environmental problems, including contamination by trace metals and petroleum hydrocarbons due to the discharges of untreated wastes and frequent oil spills (Paez-Osuna et al., 1986; Vazquez et al., 1991; Rosales-Hoz and Carranza-Edwards, 1998). Increasing trace element contamination in the Coatzacoalcos River discharge area has been related to the input of eroded soils from the catchment, promoted by the large industrial and urban development in the region, especially since the early 1980s (Ruiz-Fernández et al., 2012).

2.9. Bluefields Bay (Nicaragua, NIC)

Bluefields Bay is an estuarine lagoon system, catalogued as a Ramsar Site (#1138 since 2001). It has an extension of 176 km² and maximum depth of 4 m. Four rivers flow into the bay (Escondido, Caño Negro, Kukra, and Torsuani) although it is mainly influenced by the fluvial discharges of the Escondido River (mean yearly flow of 850 m³ s⁻¹), carrying a high sediment load (1.4 a 22.3 kg m⁻³ per day) (Brenes et al., 2007). The bay sustains an important artisanal fishing activity (shrimp, fish, oysters and clam). Siltation is a major environmental problem of the bay, which increased from 5 (1970) to 42 (1998) mm year⁻¹, attributed to socio-economic development in the area and the occurrence of natural disasters (Dumailo, 2003). Bluefields bay is affected by erosion and wastewater discharges from diverse industrial activities (e. g. mining, tannery, milk processing) and human settlements along the Escondido River watershed (RSIS, 2020a).

2.10. Almirante Bay (Panama, PAN)

The Almirante Bay is a semi-enclosed lagoon system in Bocas del Toro Archipelago, Panama. It is connected to the Caribbean Sea, and has relatively clear and oligotrophic waters, which sustain well-developed coral reefs. It has a maximum depth of 32 m, and receives fluvial inputs from several small rivers and creeks, and sediment plumes from the rivers Sixaola and Changuinola, through the oceanic inlet Boca del Drago. The presence of reef barriers and mangrove forests considerably reduce the effect of wind and meteorological events during the hurricane season. Environmental concerns are increasing untreated wastewater discharges, promoted by rapid tourism and population growth in Bocas del Toro; inland erosion and agricultural runoff due to expansion of pasture lands and plantations of banana and teak; as well as heavy metal and hydrocarbon contamination, and alteration of the sedimentation patterns caused by dredging and shipping activities through the bay (D'Croz et al., 2005; Seemann et al., 2014).

2.11. Gulf of Cariaco (Venezuela, VEN)

The Gulf of Cariaco is a marginal basin, one of the most productive in the Caribbean Sea, owing to nutrient delivery by seasonal upwelling and numerous watercourses (notably the rivers Manzanares and Carinicuao). The Gulf of Cariaco is ~60 km long and ~15 km wide, has an extension of 642 km² and a maximum water depth of 85 m (Van Daele et al., 2011). Bottom water moves in the west-east direction along the south coast until El Saco, a shallow sediment deposition zone where mangrove forests dampen the movement of water and waves. It sustains commercially important fishery resources (e.g. sardine, spadefish and shrimp). It is contaminated by trace metals and hydrocarbons, owing to industrial and domestic wastewaters transported by rivers, and intense marine traffic (Márquez et al., 2005; Romero et al., 2013). Carinicuao River flows through ~170 km of important production areas and it is in the most seismic activity zone in Venezuela.

3. Methods

The IAEA - RLA/7/012 project organized a radiochronology experts meeting to prepare a methodological guide to ensure that the collection, processing and analysis of the sediment cores used in this study followed a standardized procedure (OIEA, 2020b).

3.1. Sampling

Sediment cores were collected by using a gravity corer UWITECTM with a transparent liner (1 m long, 8.9 cm inner diameter) between 2008 and 2009 at coastal locations (Fig. 1). None of the cores presented laminations nor signs of sediment disturbance (e.g. sediment cracks or infaunal burrows). The cores were immediately transported to the counterpart laboratory, where they were extruded and subsampled at 1 cm intervals. The mass of each section was recorded before and after drying (most countries at ≤ 60 °C for at least 48 h, although in some countries samples were freeze-dried). Sediments were ground to powder by using porcelain mortar and pestle (except the aliquots used for grain size analysis). The samples were stored in polyethylene bags and shipped to previously assigned laboratories for further analysis. All concentrations and activities are expressed on a dry weight basis.

3.2. Laboratory analysis

²¹⁰Pb analyses were carried out at Servicio Académico de Fechado (ICML-UNAM, Mexico), Centro de Estudios Ambientales de Cienfuegos (CEAC-Cuba) and Radiometrics Laboratory (IAEA-EL, Monaco). ²¹⁰Pb_{tot} activities were measured through its daughter product ²¹⁰Po by alpha spectrometry, assuming secular equilibrium between the two isotopes (Ruiz-Fernández and Hillaire-Marcel., 2009). Also, gamma-ray spectrometry (HPGe well-detectors) was used to determine ²²⁶Ra (through the ²¹⁴Pb gamma emission line of 352 keV) and ¹³⁷Cs (661.5 keV). Counting uncertainties were generally ≤6% for alpha spectrometry and ≤10% for gamma spectrometry. Replicate analyses (n = 12) of the standard reference material IAEA-300 (Radionuclides in Baltic Sea sediment) confirmed good agreement (within 2σ) for ²¹⁰Pb_{tot} and ¹³⁷Cs.

Loss on ignition was determined as a rough estimation of organic matter (OM) content by combusting sediments in a muffle furnace at 550 °C (LOI₅₅₀). Grain size distribution was assessed by using a Malvern laser diffraction equipment. Aluminum concentrations were analyzed with an X-ray fluorescence (XRF) system (Spectro X lab 2000, Spectro Analytical Instruments) at IAEA-EL, Monaco. Replicate analysis (n = 6) of the reference materials IAEA-158, IAEA-405 and IAEA-433 indicated good agreement between certified and analytical values for most elements, with accuracies above 90% and uncertainties below 8% for most of the elements determined.

3.3. Data treatment

3.3.1. Statistical analysis

Basic statistics were obtained for all dataset variables, and analysis of variance (ANOVA) with Tukey post-hoc test (95% confidence) was used to evaluate differences in the mean values of these variables among the sediment cores. Cluster analysis was performed to identify which geochemical variable ($\rm LOI_{550}$, sand, silt, clay and Al) better explained the variation of radionuclide activities observed within the sampling sites.

3.3.2. ²¹⁰Pb chronology

Excess ²¹⁰Pb (²¹⁰Pb_{ex}) was determined by the difference between total ²¹⁰Pb (²¹⁰Pb_{tot}) and supported ²¹⁰Pb (²¹⁰Pb_{sup}) activities, estimated from gamma-spectrometry derived ²²⁶Ra activities. Depending on the observed profiles, ²¹⁰Pb chronologies were computed either by the constant sedimentation constant flux model (CFCS; Krishnaswamy et al., 1971) or by the constant flux model (CF; Robbins and Edgington, 1978), following the methodology and spreadsheet in Sanchez-Cabeza and Ruiz-Fernández (2012). Uncertainties were usually \leq 15% for age, MAR and SAR. The ¹³⁷Cs depth profiles were used to derive a stratigraphic marker to attempt the corroboration of the ²¹⁰Pb age models.

3.3.3. ²¹⁰Pb_{ex} and ¹³⁷Cs inventories

The radionuclide inventories were calculated by integrating over the core length the product of radionuclide activity (Bq kg⁻¹) and dry mass (kg) divided by the cross sectional area of the corer (Sanchez-Cabeza and Ruiz-Fernández, 2012). Radionuclide activities in unmeasured sections were interpolated *versus* mass depth (g cm⁻²). Uncertainty was calculated by quadratic uncertainty propagation. Differences among cores were assessed through Student's *t*-test.

4. Results

4.1. Radionuclide spatial variability

Activities ranges were 4.5–259 Bq kg⁻¹ for ²¹⁰Pb_{tot} and 4.5–50.6 Bq kg⁻¹ for ²²⁶Ra (Table 1, Fig. 2). Most cores showed ²¹⁰Pb_{tot} activities <100 Bq kg⁻¹ (except CUB, HON, MEX and VEN) and ²²⁶Ra activities < 20 Bq kg⁻¹ (except COL, MEX and VEN). ¹³⁷Cs ranged from <LD (0.1 Bq kg⁻¹) to 7 Bq kg⁻¹, with lowest values (<2 Bq kg⁻¹) observed in COL, CUB and HAI. The inventories of ²¹⁰Pb_{ex} (Bq m⁻²) ranged from 2965 ± 22 in PAN to 22,230 ± 609 in MEX, accounting for ²¹⁰Pb_{ex} fluxes between 92 ± 1 and 693 ± 19 Bq m⁻² year⁻¹, respectively; whereas the ¹³⁷Cs inventories ranged from 47 ± 2 Bq m⁻² year⁻¹ in HAI and 1,055 ± 54 Bq m⁻² year⁻¹ in MEX (Table 1, Fig. 3). A cluster analysis was performed to evaluate the main source of variability of ²¹⁰Pb_{tot}, ²²⁶Ra and ¹³⁷Cs activities in the database (Fig. 4). Results showed two main groups (LOI₅₅₀ and sand; silt, ²²⁶Ra, ²¹⁰Pb_{tot}, ¹³⁷Cs, Al and clay), and in the final partition, separate clusters were observed for i) ²¹⁰Pb_{tot}, ¹³⁷Cs and Al, and ii) ²²⁶Ra and silt.

4.2. ²¹⁰Pb chronologies and ¹³⁷Cs depth profiles

²¹⁰Pb_{tot} depth profiles (Fig. 5) showed decreasing values downcore,

but without clear exponential trends. The CF model was used to calculate all chronologies except for the MEX core, dated by using a 2 piecewise CFCS model because the presence of $a^{210}Pb_{ex}$ depleted segment, attributed to recent deposition of old sediments, excavated in the Coatzacoalcos river (Ruiz-Fernández et al., 2012). Most sediment records spanned ~100 years (Table 1S) except four younger cores COL, HON and JAM (owing to high sedimentation), as well as MEX (due to its short length).

 137 Cs activities (Fig. 5) were detectable since the bottom sections of the younger cores COL (1973), HON (1947), JAM (1940) and MEX (1956), since the mid-1940s in cores PAN and VEN, and since the early 1990s in HAI, according to the 210 Pb-derived age models (Table 1S). Cores CUB, DOM, GUA and NIC showed detectable 137 Cs in sediments older than the first nuclear weapon test (1945, Trinity test). Only PAN and GUA showed a maximum that, according to the 210 Pb-age models, could be attributable to the 137 Cs global fallout maximum (1962–1964). All cores showed higher 137 Cs activities in the top segments.

4.3. Sediment accumulation

Mass accumulation rates (MAR) and sediment accumulation rates (SAR) varied widely among and within the cores (Fig. 6). Highest MAR was usually below 1 g cm⁻² year⁻¹, and highest SAR was usually below 2 cm year⁻¹. MAR in cores COL, CUB and DOM, and SAR values in COL, CUB, DOM and JAM were significantly (p < 0.05) higher than the other cores (Fig. 6, Table 1). Both MAR and SAR showed steady increases in most cores, except CUB, DOM and HON, where clear maxima were observed at intermediate depths (up to 10 cm year⁻¹).

5. Discussion

5.1. Radionuclide spatial variability

The overall database showed an important variability of 210 Pb_{tot} and 226 Ra activities in the region (Fig. 2). Generally speaking, activities were low, with 210 Pb_{tot} activities in HON, MEX and VEN being significantly (p < 0.05) higher than in the rest of the cores, and highest values of 226 Ra were found in COL, MEX and VEN.

Activities of ²¹⁰Pb_{tot} in sediments are the result of ²¹⁰Pb fallout (in coastal sites mainly controlled by atmospheric ²¹⁰Pb concentrations and rainfall), in situ production owing to 222Rn radioactive decay, and land surface runoff (mainly caused by fluvial inputs and watershed erosion). According to a simulated global distribution of monthly mean 210 Pb concentrations in surface air, and the latitudinal variation of the annual average total ²¹⁰Pb deposition fluxes (Liu et al., 2001), WCR is within a latitudinal band characterized by low atmospheric ²¹⁰Pb fluxes, very likely owing to a low ratio of the continent land mass/ocean area. The scarce data on ²¹⁰Pb fluxes in the WCR latitudes show values ranging from 133 \pm 83 to 198 \pm 112 Bq m⁻² year⁻¹ for the latitudinal band 10–30°N (Baskaran, 2016). However, ²¹⁰Pb_{ex} deposition fluxes in most cores (Table 1; mean 210 Pb_{ex} flux of 284 ± 116 Bq m⁻² year⁻¹) fell within the common range of ²¹⁰Pb_{ex} water-sediment deposition flux in coastal areas in the northern hemisphere (100–450 Bq m^{-2} year⁻¹; Preiss et al., 1996). These relatively high values were caused by the high contribution of erosion, mainly through river inputs, and some degree of fine sediment focusing at the sampling sites. The highest ²¹⁰Pb_{ex} fluxes were observed in HON and MEX (556 \pm 5 and 693 \pm 19 Bq m⁻² year⁻¹, respectively), which were collected in coastal environments surrounded by extensive catchment areas and irrigated by large rivers: 21,336 km², 910 m³ s⁻¹ for the Coatzacoalcos River, Mexico (CONAGUA, 2018), and 3757 km², 370 m³ s $^{-1}$ for Chamelecon River, Honduras (USGS, 2002; Seeliger and Kjerfve, 2000).

 226 Ra in soils is mostly associated to the natural abundance of 238 U of the parent rock mineralogical composition. During rock weathering (leaching and dissolution of minerals), 226 Ra can be solubilized in surface or ground waters, and transferred with the water flow until it

Table 1

Basic statistics of radionuclide data, Al concentrations, accumulation rates, ¹³⁷Cs inventories and ²¹⁰Pb_{ex} inventories and fluxes in sediment cores from the Wider Caribbean Region.

Variable						Core					
²¹⁰ Pb _{tot} (Bq kg ⁻	COL 1)	CUB	DOM	GUA	HAI	HON	JAM	MEX	NIC	PAN	VEN
N	36	35	47	18	14	37	31	26	48	53	60
Min	18.2	11.0	4.5	10.1	6.5	53.1	14.6	39.7	10.7	7.1	23.6
Max	54.7	259.0	56.5	92.1	40.5	184.5	49.2	114.0	61.8	69.6	201.0
Mean	36.5	41.3	21.4	44.4	26.0	134.7	32.3	67.0	39.4	37.8	96.3
SD	11.7	54.7	15.6	30.6	10.8	29.7	11.1	17.2	12.9	20.6	37.2
Median	35.9	24.7	14.5	44.8	29.8	131.2	37.6	64.5	41.2	38.1	96.7
226 Ra (Bu kg ⁻¹)											
N N	36	35	47	18	15	37	31	26	48	11	60
Min	145	51	45	10.4	60	117	80	13.9	10.2	6.8	23.9
Max	28.1	14.1	7.6	13.1	11.6	11.7	15.0	27.0	14.4	8.0	50.6
Mean	20.1	95	5.0	11.0	0.1	11.7	10.0	20.4	19.1	7.4	40.6
SD	4.3	2.4	0.7	0.7	1.5	0.0	12.1	20.4	12.1	0.4	5 3
Modion	7.5	2.4	0.7 E 0	12.0	1.5	11.7	1.5	21.2	1.0	75	20.0
Wedian	23.3	0.1	3.9	12.0	9.0	11./	12.4	21.2	12.2	7.5	39.9
²¹⁰ Pb _{ex} (Bq kg ⁻¹)											
N	36	8	36	12	7	36	30	26	45	49	23
Min	1.6	5.1	0.3	12.8	10.2	41.4	1.7	26.9	3.7	5.3	91.5
Max	30.2	19.1	50.2	86.1	28.7	176.6	37.5	100.8	39.7	58.1	193.4
Mean	14.4	9.4	18.9	55.4	19.8	127.0	21.0	54.1	24.5	30.4	123.4
SD	8.4	4.3	15.4	23.9	5.8	29.6	11.4	17.4	9.9	17.1	25.0
Median	13.7	8.0	17.9	60.1	18.5	122.7	26.6	53.0	24.5	30.2	117.4
Son Cs (Bq kg -)	07	01	47	10	15	07	01	26	40	20	20
N Mire	2/	21	4/	18	15	3/	31	20	48	20	20
NIIII	0.9	0.3	1.0	7.0		2.4		1.4	0.8		1.4
Max	1.9	1.0	3.0	7.0	1.1	4./	4.1	3.9	3.8	2.6	3.5
Mean	1.2	1.1	1.9	4.5	0.4	3.8	2.1	2.4	2.5	1.6	2.4
SD	0.3	0.4	0.5	1.6	0.5	0.6	1.0	0.8	0.8	0.5	0.6
Median	1.2	1.1	1.8	5.0	*	3.9	1.8	2.3	2.7	1.6	2.4
Al (%)											
Ν	60	29	65	36	48	37	18	26	42	55	53
Min	1.1	4.7	4.2	4.3	1.3	8.6	2.8	4.5	6.0	3.5	8.9
Max	6.1	6.1	5.4	5.4	2.3	9.1	5.1	7.9	10.4	4.4	11.1
Mean	3.5	5.8	4.9	4.8	1.9	8.9	3.7	6.2	9.4	3.9	10.3
SD	1.2	0.3	0.3	0.3	0.3	0.1	0.8	0.9	0.8	0.2	0.5
Median	3.7	5.9	4.9	4.7	1.9	8.8	3.3	6.2	9.6	3.9	10.3
MAR (g cm ⁻² y	ear ⁻¹)									10	
N	65	22	65	28	45	36	110	26	45	49	45
Min	0.6	0.4	0.03	0.03	0.03	0.3	0.3	0.7	0.03	0.01	0.1
Max	1.1	8.0	9.9	0.2	0.4	0.9	0.6	0.8	0.4	0.2	0.5
Mean	0.8	2.1	1.6	0.1	0.2	0.5	0.4	0.7	0.2	0.1	0.3
SD	0.1	1.6	2.3	0.1	0.1	0.2	0.1	0.1	0.1	0.0	0.1
Median	0.8	2.1	0.6	0.1	0.3	0.4	0.4	0.7	0.2	0.1	0.3
SAR (cm year ⁻¹)											
Ν	65	22	65	28	45	36	110	26	45	49	45
Min	1.3	0.6	0.1	0.1	0.04	0.3	0.9	0.4	0.04	0.04	0.1
Max	2.9	10.3	11.3	1.1	2.0	1.4	6.2	0.6	1.5	2.8	1.7
Mean	1.9	2.5	1.8	0.3	0.5	0.7	1.9	0.5	0.5	0.5	0.7
SD	0.3	2.0	2.3	0.2	0.4	0.3	1.1	0.1	0.4	0.5	0.4
Median	1.8	2.3	0.8	0.3	0.4	0.6	1.4	0.6	0.4	0.4	0.6
Radionuclide inventories (Bq m ⁻²)											
²¹⁰ Pb _{ex}	5807	7085	8908	5260	3890	17838	4621	22230	7302	2965	14328
±	148	214	257	37	89	169	51	610	195	22	406
¹³⁷ Cs	360	169	1056	390	47	1050	823	905	615	100	393
±	12	13	28	9	4	27	13	54	14	2	18
Elux (Ba m^{-2} vear ⁻¹)											
²¹⁰ Ph	181	221	278	164	121	556	144	693	228	92	447
+ Dex	5	7	8	1	3	5	1 T T 2	19	6	1	13
<u> </u>	5	,	5	±	5	5	-	1.7	5	1	13

 $\rm NA=not$ available; *below minimum detectable activity (0.1 Bq $\rm kg^{-1}).$



Fig. 2. Boxplot of radionuclide activities, Al concentrations and sedimentation rates in sediment cores collected from the Wider Caribbean Region.

becomes adsorbed onto suspended or settled particles. ²²⁶Ra sorption is most effective by smaller grains, owing to a greater surface to volume ratio (Webster et al., 1995), by clays, metal hydroxides and humic substances, and under conditions of neutral to alkaline pH (Urso et al., 2019) and low salinities (IAEA, 1990). ²²⁶Ra can also be enriched in soils owing to continued and long-term application of phosphate fertilizers (Khan et al., 2018). ²²⁶Ra in marine sediments is mostly derived from local production by ²³⁰ Th day plus the fluvial contribution of soil particles.

The low ¹³⁷Cs activities are the result of low fallout in lower latitudes of the northern hemisphere (Aoyama et al., 2006), the time elapsed since ¹³⁷Cs fallout (\sim 1963) and the high solubility of ¹³⁷Cs in seawater.

According to UNSCEAR (2000), the 10° – 20° N latitudinal band (where most of the sampling sites are located; Fig. 1) received ~11% of the total radionuclide deposition on the northern hemisphere from atmospheric nuclear testing. In addition, only ~65% the ¹³⁷Cs deposited in 1963 would remain in the environment, owing to radioactive decay, since no significant ¹³⁷Cs has been released in the region since 1960s. In addition, high ¹³⁷Cs solubility in seawater causes that >99% of ¹³⁷Cs in marine waters remains in solution (Molero et al., 1995 and references therein; Atwood, 2013) and, owing to ocean stratification, ¹³⁷Cs is not efficiently transported to bottom waters. In addition, ¹³⁷Cs can be transferred by runoff to the coastal zone were, owing to salinity changes, can be recycled and scavenged mostly by fine particles (e.g. organic matter and



Fig. 3. Inventories of excess ²¹⁰Pb and ¹³⁷Cs in sediment cores from the Wider Caribbean Region.



Fig. 4. Dendrogram from a cluster analysis of radionuclide activities and some geochemical variables in sediment cores from the Wider Caribbean Region.

clay minerals) to be finally deposited to the sediments.

Data on ¹³⁷Cs inventories in shallow coastal areas, influenced by fluvial inputs, in similar latitudes are scarce in the literature. The ¹³⁷Cs inventories (47 \pm 2–1056 \pm 54 Bq m⁻²) were low in comparison with values reported for areas influenced by river discharges in the Adriatic Sea (up to 30,000 Bq m⁻²; Papucci et al., 1996), where prodelta sediments seem to have a high trapping efficiency for fallout ¹³⁷Cs, but receive additional ¹³⁷Cs inputs through surface soils, eroded from the catchment, that were previously contaminated by ¹³⁷Cs fallout. The mean ¹³⁷Cs inventory in this study (537 \pm 221 Bq m⁻²) fell within the range reported for shallow nearshore sediments in NW Atlantic (144–544 Bq m⁻²; Livingston and Bowen, 1979). The highest ¹³⁷Cs inventories observed in DOM, HON and MEX were comparable to values reported for Sagua Estuary (1160 \pm 162 Bq m⁻²) and undisturbed pasture fields near Cienfuegos Bay in Cuba (1200 \pm 198 Bq m⁻², Díaz-Asencio et al., 2009 and references therein).

Cluster analysis showed that the main carrying phase of $^{210}Pb_{tot}$, ^{226}Ra and ^{137}Cs to sediments are fine grained particles. The clustering of $^{210}Pb_{tot}$ and ^{137}Cs activities with Al concentrations (as indicator of terrigenous input) and ^{226}Ra activities with silt-sized particles suggests that the main variability source of radionuclides is fluvial runoff to the study sites. Aluminosilicate particles are considered a primary phase carrying $^{210}Pb_{tot}$ from the water column to the sea floor in tropical areas (Legeleux et al., 1996) and it seems that, regionally speaking, ^{226}Ra is contained in minerals predominantly present in medium-sized (silt) particles.

5.2. ²¹⁰Pb chronologies and ¹³⁷Cs depth profiles

Most ²¹⁰Pb_{tot} depth profiles showed departures from the exponential decreasing trend, and these non-monotonic profiles could be indicative of post-depositional sediment mixing, since physical stirring due to strong erosive currents produced by winds and tides, as well as bioturbation, often result in obliterating the original ²¹⁰Pb signal. However, other processes can have similar effects on sediment records; for

instance, reduced ²¹⁰Pb_{ex} activities at a core surface segment (e.g. cores GUA, JAM, NIC, PAN) could be the result of very rapid sedimentation (Appleby, 1998). As shown by Krishnaswamy et al. (1971), ²¹⁰Pb_{ex} activities in the sediments result from the balance between the ²¹⁰Pb_{ex} flux to the sediments and the sediment load (i.e. mass accumulation rate, MAR); this implies that ²¹⁰Pb_{ex} activities would decrease as a result of dilution by a higher MAR, or increase if the MAR was lower.

²¹⁰Pb_{ex} profiles mostly shaped by radioactive decay would be only expected in environments where steady state sedimentation conditions prevail, but such conditions are hardly met in the coastal zone. Firstly, because of the multiplicity of particle sources and the highly variable sedimentation conditions that characterize the estuarine areas (UNESCO, 1978); and secondly, owing to the "Great Acceleration" of population growth and industrialization (Steffen et al., 2007) that characterizes the Anthropocene (i.e. from 1950 onwards; AWG, 2019). One of the most conspicuous evidences of global change is the increasing erosion owing to land use changes (Vitousek, 1994) and the increment in sediment accumulation rates in the coastal environments. Consequently, ²¹⁰Pb_{ex} profiles from anthropized areas in the dynamic coastal environment will seldom show an exponential decay trend. In this study, the deviations observed in the ${}^{210}P\dot{D}_{ex}$ profiles were interpreted as changes of sediment accumulation, under the assumption that post-depositional mixing was likely minor.

Independent validation of the chronology is essential for a high level of confidence in the dating results, for instance by using ¹³⁷Cs activity profiles as time stratigraphic markers (Smith, 2001). However, the ¹³⁷Cs stratigraphic profiles were of limited usefulness to corroborate ²¹⁰Pb-derived age models, owing to the lack of a clear ¹³⁷Cs activity maximum (except PAN and GUA). Other processes that might limit the effective use of the ¹³⁷Cs as stratigraphic marker in the coastal zone are the input by runoff and postdepositional remobilization. Fluvial input of ¹³⁷Cs-bearing soils eroded from the catchment could mask the expected global fallout maximum of ¹³⁷Cs in ~1963 and could also explain the relatively high ¹³⁷Cs activities observed in the recent core sections, in comparison to the older ones. Mobility problems for ¹³⁷Cs have also been recurrently reported since long ago (e.g. Davis et al., 1984; Carpenter, 1997; Klaminder et al., 2012; Wang et al., 2017). Some sediment components (clay minerals, organic matter, carbonates and iron-manganese oxy-hydroxides) are known to bind ¹³⁷Cs (Ashraf et al., 2014) and can be affected by burial diagenesis, which would promote the release and downward transport of ¹³⁷Cs in the pore waters through the sediment column, which could explain the presence of ¹³⁷Cs in layers presumably deposited before the onset of nuclear tests (such as in cores CUB, DOM, GUA and NIC). Higher salinities in the overlying water column and pore waters can produce ¹³⁷Cs desorption and upward migration, due to increasing competing ions such as sodium and potassium, whereas lower in the sediments, as the release of NH⁴⁺, Fe²⁺ and Mn⁺² in anaerobic conditions promotes the ion-exchange displacement of ¹³⁷Cs from sediments (Evans et al., 1983). Despite all these limitations, ¹³⁷Cs maxima were compatible with the chronologies obtained for PAN and GUA; and in addition, cores COL, HON, MEX and



Fig. 5. Radionuclide activity (total ²¹⁰Pb, red; ²²⁶Ra, blue; ¹³⁷Cs, yellow) depth profiles (mass depth) in sediment cores from the Wider Caribbean Region. (For interpretation of the references to colour in this figure legend, the reader is referred to the Web version of this article.)



Fig. 6. Temporal variation of sediment accumulation rates (SAR) and mass accumulation rates (MAR) in sediment cores from the Wider Caribbean Region.

VEN showed ¹³⁷Cs activities in all sections, indicating that these sediments accumulated after the onset of the nuclear testing period, in agreement with the ²¹⁰Pb-derived dates.

Owing to all the problems to corroborate ²¹⁰Pb-derived chronologies with ¹³⁷Cs depth profiles, other stratigraphic markers are needed for low latitude coastal environments. Plutonium isotopes have lower postdepositional mobility than ¹³⁷Cs, for which their use as stratigraphic markers might be a good alternative (Hancock et al., 2011). Plutonium isotope depth profiles have been used to corroborate ²¹⁰Pb chronologies in coastal areas of Cuba and Mexico, including cores CUB (Diaz-Asencio et al., 2011; Corcho Alvarado et al., 2014) and MEX (Ruiz-Fernández et al., 2012). Furthermore, independent chronostratigraphic features in pollen, diatom or anthropogenic substances (e.g. heavy metals) might be used as time markers in recent sediments (UNESCO, 1978; Appleby, 1998). The cores COL, CUB and MEX have been previously used to study trace element contamination. High mercury concentrations recorded in core COL between 1970s and 1980s were related to the discharge of mercury contaminated wastes to Cartagena Bay from a chlor-alkali plant that used mercury as a cathode in the production of chlorine, and the maximum values in the late 1970s, and the subsequent decreasing concentrations, were related to the plant closure (Orani et al., 2020). In core CUB, the temporal variability of the anthropogenic fluxes of Pb, Zn, Sn, Hg and Cr accurately described the increasing impact from the urban and industrial activities around Havana Bay between the beginning of past century and 1980s; and the drastic reduction of these fluxes since 1990s was attributed to the combination of coastal zone management policies and the economic contraction of Cuba (Díaz-Asencio et al., 2011). In core MEX, concentrations of major (Al, Ca) and trace (Sr, Rb, As, Cu, Ni, V and Zn) elements showed a change of sediment and contamination sources since early 1980s, resulting from the large industrial and urban development in the Coatzacoalcos River discharge area, indicating that the increased contaminant loads were related to eroded soils from the catchment, due to extensive land use changes (Ruiz-Fernández et al., 2012). The data on plutonium isotope activities

and heavy metals concentrations from these works can be found in the original sources.

In this work, efforts were made to corroborate the age models by combining the ¹³⁷Cs data with ²¹⁰Pb_{tot} stratigraphic features (Fig. 5) and MAR profiles (Fig. 6), based on historical data available at each site. Owing to the influence of changes in sediment density (due to compaction with time) SAR values must be taken with caution; MAR was preferred over SAR to explain these historical variations. ²¹⁰Pb chronologies proposed for most cores are only preliminary and may be adjusted in the future if additional information is available; further research on contamination in the cores might contribute to this purpose.

5.3. Colombia

MAR at Cartagena Bay was moderately variable (Table 1, Fig. 6). The main source of variability is apparently related to interventions in the Dique channel, through which most sediments enter the bay. The lowest values were recorded before early 1980s and the increments observed afterwards were attributed to the last rectification of the Dique channel (1975–1983), which increased 26% the water flow, and thus sediment delivery (UNC, 2007).

5.4. Cuba

MAR values were highly variable between 1910s and 1980s (Fig. 6), with maxima associated with the occurrence of extreme meteorological events in 1926 (the Great Hurricane), 1944 (Cuba–Florida hurricane) and 1982 (the heaviest rains reported in Havana over the 20th century). Since 1990, MAR values are similar to those observed in 1900, which were attributed to the reduction of economic activities in the entire catchment area between 1990 and 1995 (associated with the economic contraction of Cuba in the early 1990s due to the economic collapse of the socialist block), and with environmental measures implemented since 1998 to reduce the input of suspended solids to Havana Bay, which

have contributed to restore natural sediment flows in this coastal ecosystem (Díaz-Asencio et al., 2011).

5.5. Dominican Republic

MAR values where almost constant (within the uncertainties) through the sediment record, and maxima are in close agreement with the occurrence of extreme events (e.g. San Cristobal storm in 1945; hurricane David and storm Frederick in 1978; Georges hurricane in 1998). The lower reach of Haina River is impacted by urban and industrial activities, sand extraction, deforestation and oil residues (Gómez Mena et al., 2008). However, excluding MAR maxima, the MAR range (from <0.06–0.8 cm year⁻¹) would place the port of Haina River among the areas least impacted by siltation in this study.

5.6. Guatemala

The MAR profile showed almost constant values between 1900s and early 1930s, a steady increase between early 1940s and 2000s (~2 fold within the period), which speeded up through the last decade of the record (~4 fold with respect to the values recorded in early 1940s). MAR increments were attributed to deforestation trends in the catchment area of Amatique Bay. The country has lost ~11% of its forest cover within 1991/93–2001 owing to the development of agriculture and cattle raising activities, the growth of disordered human settlements, and the occurrence of prolonged drought periods and forest fires (IARNA, 2006).

5.7. Haiti

In the Bay of Port-au-Prince, MAR values showed a steady increase throughout the record, albeit at a faster pace since 1970s. Deforestation is a pervasive environmental problem in Haiti since the beginning of the 19th century. Forest cover in Haiti has decreased from 60% in 1923 (Johnson Williams, 2011) to 30% in 1940, 10% in 1970 and to 4% in 2010 (FAO, 2010), and the primary forest has declined from 4.4% in 1988 to 0.32% in 2016 (Hedges et al., 2018). Accelerating deforestation and soil erosion since early 1950s is attributed to environmentally unsound agricultural practices, rapid population growth; and increasing logging operations since 1970s, owing to Port-au-Prince's intensified demand for charcoal (Johnson Williams, 2011).

5.8. Honduras

MAR values in HON were almost constant (within the uncertainties) along the core, and are relatively high in comparison with the rest of the cores. Forest cover loss in Honduras has been reported to be very high (\sim 23% between 1958 and 1990, and \sim 30% between 1990 and 2005; FAO, 1960 and 2010) owing to extensive agriculture and logging activities and the development of human settlements. No remarkable feature in the MAR profile was observed, that could serve as a time marker to corroborate the ²¹⁰Pb-derived age model.

5.9. Jamaica

MAR values were almost constant and relatively high (0.58 \pm 0.02 g cm $^{-2}$ year $^{-1}$) between 1940 and the early 1960s, values declined sharply between 1960s and 1970s and were constant until middle 1990s, since when it showed a steady but rapid increase to reach a maximum comparable to values observed at the beginning of the record (0.59 \pm 0.04 g cm $^{-2}$ year $^{-1}$) in 2008. The declining MAR values recorded between 1960 and 1970 are most likely related with the construction of the Portmore Causeway (1968–1969), which was built to link Kingston with the Portmore development; however, before development began, about 800 ha of lowlands were filled by dredging the harbor, and a 19 km dyke structure was constructed to prevent flooding of the low-lying area from the Rio Cobre.

Environmental degradation of Kingston Bay has been recognized since late 1960s; the main impacts being eutrophication, sedimentation, and solid waste build-up, mostly related to urbanization and agricultural runoff. The harbor receives fresh water from two main rivers, a network of 19 major gullies that traverse the city and effluents from sewage treatment facilities. Sand mining, deforestation and poor farming practices lead to soil erosion in the river watersheds, promoting siltation; gullies carry solid wastes from residential and industrial areas, and sewage is the major cause of total suspended solids (Webber et al., 2019). Highest MAR values since the late 1990s might be related to i) the construction of Greater Portmore in the mid-1990s, which promoted a large population growth, ii) increasing upland erosion, as a result of large-scale agriculture since 1980s (mainly coffee plantations, FAO, 2003 and 2008), and iii) the occurrence of extreme events, which promote high sediment discharges, such as a flash flood in 1998, a flood in 2002, the hurricane Dean in 2007, and the tropical storms Nicole and Gustav in 2008. All these events are well represented as maxima in the MAR profile.

5.10. Mexico

The sediment record from the Coatzacoalcos river estuary spanned 52 years. The ²¹⁰Pb_{ex} profile showed a strong discontinuity between 9 and 14 cm depth, which was attributed to the accumulation of old sediments (depleted in ²¹⁰Pb_{ex}). According to the ²¹⁰Pb-derived age model, the core section at the base of this discontinuity corresponded to 1984 ± 6, and it was related to deep excavations (2 m diameter and up to 13m depth) in the riverbed during the construction of a bridge at the lower reach of the Coatzacoalcos River in 1980, in agreement with the ²¹⁰Pb chronology (Ruiz-Fernández et al., 2012). Despite extensive land use changes and industrial development in the study area along the period comprised in the sediment record, the mean MAR values below (0.80 ± 0.08 cm year⁻¹, 1956–1982) and above (0.69 ± 0.16 cm year⁻¹, 1994–2008) the discontinuity were comparable, likely due to the high flow rate of the Coatzacoalcos River, which washes out and redistribute the eroded soils transported from the catchment.

5.11. Nicaragua

MAR values progressively increased over time, although the pace observed since the late 1970s was higher than in the previous decades. The small MAR maximum in 1987 \pm 1 was attributed to higher sedimentation loads promoted by rainfall occurred during hurricane Joan in 1988. One of the main environmental problems in Bluefields Bay, increased siltation attributed to the land erosion from the catchment of Escondido River, is clearly observed in the MAR profile, especially since late 1970s. Similar findings were reported from the study of a sediment core collected in the lower reach of Escondido River (Dumailo, 2003), the main tributary of Bluefields Lagoon. The increasing accumulation rates were associated to deforestation in the Escondido River catchment, owing to extensive cattle raising, expansion of the agricultural frontier, recurrent forest fires and the proliferation of dirt roads.

5.12. Panama

MAR values steadily increased over time in Almirante Bay, most likely as a result of population growth and consequent human impacts, at a low pace since the beginning of the 20th century to 1970s, when agriculture was the most important economic activity in the region, and at higher pace since 1980s, owing to the fast development of tourism activities. The MAR maximum observed at the end of the profile was related to the occurrence of an intense rainfall occurred in November 2008.

5.13. Venezuela

MAR values showed a steady increase since the beginning of the record, with episodic maxima (e.g. 1960, 1980) and minima (1988 and 2002) for which we do not have a clear explanation. The continuous MAR increments can be explained by the periods i) 1940–1970s, with the development of the sardine industry, ii) 1970s–2000s, as after the nationalization of the oil industry, extensive country-wide LUC took place, and iii) a steeper pace since 2000, with the reactivation of the cocoa cultivation and export activities, which currently accounts for the 90% of Venezuelan exports. A similar MAR temporal profile was reported by Yang et al. (2016) for Jiaozhou Bay, China, where the temporal sedimentation dynamics was explained to be driven by changes in economic policies, in response to population and industry expansion.

6. Concluding remark

Although with different intensity, most sediment cores showed the impact of global change through the Anthropocene (since middle 1950s), evidencing increasing sediment and mass accumulation rates with time. Despite MAR ranges in cores GUA, HAI, NIC, PAN and VEN were among the lowest in this study (Table 1), it was remarkable that the values observed within the last decade implied between ~8 and ~21 fold increments in comparison with values recorded at the beginning of the century. Increasing rates were mostly attributed to continental erosion, likely associated with LUC in the catchment areas, for reasons specific to national development realities.

7. Conclusions

A survey on temporal variations of sediment accumulation within the past 100 years was performed in 11 countries of the Wider Caribbean Region. This information is the first-ever comprehensive study on this subject in the region. Although ²¹⁰Pb dating is a powerful tool to provide retrospective information on environmental changes in coastal environments, this study confirmed that corroboration of ²¹⁰Pb chronologies with ¹³⁷Cs was difficult, owing to the input of land eroded ¹³⁷Cs enriched sediments, and to the low activities observed, which were caused by low ¹³⁷Cs global fallout in the region, high solubility of ¹³⁷Cs in seawater and radioactivity decay since the period of nuclear atmospheric testing. Thus, it is important to find alternative ways to corroborate the ²¹⁰Pb age models, including other radioactive methods, or historical information that could explain features of either ²¹⁰Pb profiles or derived mass accumulation rates, as shown in this survey.

Siltation is a common problem in all studied sites, and sedimentation rates have been steadily rising since the beginning of 20th century because of deforestation and soil erosion in the surrounding catchment areas, and poor management of urban and industrial wastes. Owing to the different characteristics of the study sites (from coastal lagoons to open waters), accumulation rates cannot be extrapolated and must be obtained for each area of interest. Retrospective studies based on ²¹⁰Pb dating are useful not only to evaluate the historical changes occurred in the coastal environments, but also to verify the usefulness of management programs to control the land-derived erosion and improve siltation problems, such as in Havana Bay (Cuba).

The results of this research must be taken with caution because (a) they stem from a single core per study site, which may not depict the conditions of the whole ecosystem; and (b) due to the lack of unequivocal validation of some ²¹⁰Pb chronologies. However, they are valuable because of the scarcity of environmental data at some of these sites, and lead to important conclusions regarding the coastal ecosystems in WCR. It is certainly expected that these preliminary results will promote further research on the interactions between anthropogenic activities and coastal environments in this region. Owing to the detrimental potential of siltation to the valuable coastal resources, it is strongly recommended to follow up on these retrospective evaluations, and to

extend them to other areas, to assist integrated coastal zone management programs in the region and elsewhere.

Declaration of competing interest

The authors declare that they have no known competing financial interests or personal relationships that could have appeared to influence the work reported in this paper.

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Appendix A. Supplementary data

Supplementary data to this article can be found online at https://doi.org/10.1016/j.jenvrad.2020.106366.

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